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Abstract: A rapid batch extraction method was evaluated to estimate potential for total dissolved solids (TDS) release by 65 samples of rock from coal and gas-bearing strata of the Appalachian Basin in eastern USA. Three different extractant solutions were considered: deionized water (DI), DI equilibrated with 10% CO2 atmosphere (DI+CO2), or 30% H2O2 under 10% CO2 (H2O2+CO2). In all extractions, 10 g of pulverized rock (<0.5-mm) were mixed with 20 mL of extractant solution and shaken for 4 hours at 50 rpm and 20-22oC. The 65 rock samples were classified as coal (n=3), overburden $(n=17)$, coal refuse that had weathered in the field $(n=14)$, unleached coal refuse that had oxidized during indoor storage (n=20), gas-bearing shale (n=10), and pyrite (n=1). Extracts were analyzed for specific conductance (SC), TDS, pH, and major and trace elements, and subsequently speciated to determine ionic contributions to SC. The pH of extractant blanks decreased in the order DI (6.0), DI+CO2 (5.1), and H2O2+CO2 (2.6). The DI extractant was effective for mobilizing soluble SO4 and Cl salts. The DI+CO2 extractant increased weathering of carbonates and resulted in equivalent or greater TDS than the DI leach of same material. The H2O2+CO2 extractant increased weathering of sulfides (and carbonates) and resulted in greatest TDS production and lowest pH values. Of the 65 samples, 19 had leachate chemistry data from previous column experiments and 35 were paired to 10 field sites with leachate chemistry data. When accounting for the water-to-rock ratio, TDS from DI and DI+CO2 extractions were correlated to TDS from column experiments while TDS from H2O2+CO2 extractions was not. In contrast to column experiments, field SC was better correlated to SC measured from H2O2+CO2 extractions versus DI extractions. The field SC and SC from H2O2+CO2 extractions were statistically indistinguishable for 7 of 9 paired data sets while SC from DI extractions underestimated field SC in 5 of 9 cases. Upscaling comparisons suggest that (1) weathering reactions in the field are more aggressive than DI water or synthetic rainwater extractants used in batch or column tests, and (2) a batch extraction method utilizing 30% H2O2 (which is mildly acidic without CO2 enrichment)

could be effective for identifying rocks that will release high amounts of TDS.

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Dear Editor,

We are pleased to submit the manuscript "Batch Extraction Method to Estimate Total Dissolved Solids (TDS) Release from Coal Refuse and Overburden" by Luis Castillo Meza, Charles Cravotta, Travis Tasker, Nat Warner, Lee Daniels, Zenah Orndorf, Tim Bergstresser, May Douglass, George Kimble, Joelle Streczywilk, Chris Barton, Stephanie Fulton, Aaron Thompson, and myself to *Applied Geochemistry*. I am the corresponding author and all my contact information is included below. Through the online submission portal, we input contact information for five suggested reviewers.

Our manuscript demonstrates that a relatively simple and rapid batch extraction test can be used to estimate the mass of TDS released from freshly unearthed rock. These results are significant and highly applied because they provide a tool to manage water quality impacts associated with any earth-moving operation (e.g., surface mining, highway construction, land development). This manuscript presents the most comprehensive data set that we are aware of that combines batch extraction data for 65 samples of rock from coal and gas-bearing strata of the Appalachian Basin in the eastern United States (all newly reported), with data from 'up-scaled' results from column leaching experiments (most previously reported) and from field monitoring stations (most newly reported). We demonstrate that a batch extraction method utilizing 30% H2O2 (which is mildly acidic) could be effective for identifying rocks that will release high amounts of TDS.

We strongly believe this work would be of great interest to the readers of *Applied Geochemistry*.

Sincerely,

Willm D Bwzyne

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Abstract

 A rapid batch extraction method was evaluated to estimate potential for total dissolved solids (TDS) release by 65 samples of rock from coal and gas-bearing strata of the Appalachian Basin in eastern USA. Three different extractant solutions were considered: deionized water (DI), DI equilibrated with 27 10% CO₂ atmosphere (DI+CO₂), or 30% H₂O₂ under 10% CO₂ (H₂O₂+CO₂). In all extractions, 10 g of pulverized rock (<0.5-mm) were mixed with 20 mL of extractant solution and shaken for 4 hours at 50 29 rpm and 20-22 °C. The 65 rock samples were classified as coal $(n=3)$, overburden $(n=17)$, coal refuse that 30 had weathered in the field (n=14), unleached coal refuse that had oxidized during indoor storage (n=20), gas-bearing shale (n=10), and pyrite (n=1). Extracts were analyzed for specific conductance (SC), TDS, pH, and major and trace elements, and subsequently speciated to determine ionic contributions to SC. 33 The pH of extractant blanks decreased in the order DI (6.0), $DI+CO₂$ (5.1), and H₂O₂+CO₂ (2.6). 34 The DI extractant was effective for mobilizing soluble SO₄ and Cl salts. The DI+CO₂ extractant increased weathering of carbonates and resulted in equivalent or greater TDS than the DI leach 36 of same material. The $H_2O_2+CO_2$ extractant increased weathering of sulfides (and carbonates) 37 and resulted in greatest TDS production and lowest pH values. Of the 65 samples, 19 had leachate chemistry data from previous column experiments and 35 were paired to 10 field sites with leachate 39 chemistry data. When accounting for the water-to-rock ratio, TDS from DI and $DI+CO₂$ extractions were 40 correlated to TDS from column experiments while TDS from $H_2O_2+CO_2$ extractions was not. In contrast 41 to column experiments, field SC was better correlated to SC measured from $H_2O_2+CO_2$ extractions versus 42 DI extractions. The field SC and SC from $H_2O_2+CO_2$ extractions were statistically indistinguishable for 7 of 9 paired data sets while SC from DI extractions underestimated field SC in 5 of 9 cases. Upscaling comparisons suggest that (1) weathering reactions in the field are more aggressive than DI water or synthetic rainwater extractants used in batch or column tests, and (2) a batch extraction method utilizing 46 30% H_2O_2 (which is mildly acidic without CO₂ enrichment) could be effective for identifying rocks that will release high amounts of TDS.

1. Introduction

 High salinity in streams downgradient of coal-mining and processing facilities in the eastern U.S. has caused fish kills and harmed sensitive aquatic organisms within the past decade (Barrett, 2015; Cormier et al., 2013a; Cormier et al., 2013b; Pond et al., 2008). At the same time, higher than normal concentrations of total dissolved solids (TDS), chloride, and bromide have been documented in the Allegheny and Monongahela Rivers in western Pennsylvania (Wang, 2014; Ziemkiewicz, 2015a), while a gradual increase in salinity attributed to chloride has been documented in major rivers in the northeastern U.S. (Kaushal et al., 2005, 2018). Such observations for coal-mine drainage in the northern Appalachian Basin may be explained by recent changes in resource extraction activities that can be influenced by residual brine in the rock, including the underground mining of coal into progressively deeper zones, the development of coal-bed methane, and the development of shale gas reserves in strata below the coal- bearing formations, notably the Marcellus Shale (Cravotta and Brady, 2015; Donovan and Leavitt, 2004; Donovan et al., 2015; Ziemkiewicz, 2015a). An understanding of the potential sources of salinity at local and watershed scales is necessary for the development of effective strategies to minimize and mitigate aquatic impacts from elevated TDS.

 Accelerated mineral weathering generally accounts for increased TDS release from coal-mining landscapes (Brady et al., 1998; Timpano et al., 2010, 2015). Although acidic drainage and TDS release are commonly attributed to abandoned mines, the use of overburden materials as topsoil substitutes and the placement of carbonate-bearing overburden materials can contribute to elevated TDS (Bernhardt et al., 2012; Cormier et al., 2013; Zipper et al., 2015). Topsoil substitution with overburden is explicitly allowed in the Surface Mining Control and Reclamation Act (SMCRA), and the application of acid-base accounting (ABA) procedures guides placement of alkaline strata with the explicit goal of decreasing acidity from sulfide oxidation (Skousen et al., 2002). Oxidative dissolution of sulfide minerals will release dilute sulfuric acid, iron, and other metal(loids). Subsequent neutralization of sulfuric acid by carbonate minerals, used in ABA to balance acid generation, results in the release of calcium, magnesium, and bicarbonate. Although silicate mineral weathering rates are slower than those of carbonates, silicates

 predominate in coal overburden and can be significant sources of calcium, magnesium, sodium, potassium, aluminum, and silicon (Brady et al., 1998; Clark et al., 2018; Hammarstrom et al., 2009). Dissolution of salts or *in situ* brines retained in the rock also releases sodium, calcium, sulfate, and chloride (as well as, barium, strontium, and bromide).

 Laboratory-scale column and mesocosm-scale lysimeter experiments have been used to predict TDS release from coal industry materials (overburden, refuse, combustion byproducts) for nearly three decades (e.g., Brady et al., 1998; Clark et al., 2018; Daniels et al., 2014a; Daniels et al., 2016; Daniels et al., 2014b; Hornberger et al., 2004; Orndorff et al., 2015). These studies have shown that (1) rock type 82 and extent of weathering influence TDS release, (2) shales and mudstones release more TDS than sandstones, and (3) unweathered rocks release more TDS than weathered materials. The specific conductance (SC) of column leachates typically starts near peak values, decreases during the first few pore volume leach cycles, and then stabilizes over the remaining leach cycles. For weathered overburden 86 materials, peak leachate SC was often less than 500 μ S cm⁻¹, a proposed regulatory limit (Cormier et al., 2013). Coal refuse produced during coal processing generated considerably higher peak SC and associated concentrations of TDS, acidity, and major and trace elements compared to overburden and interburden strata removed during mining operations (Cravotta and Brady, 2015; Daniels et al., 2014b; 90 Orndorff et al., 2015). In one study (Daniels et al., 2014b), columns (0.0012 m^3 rock) were upscaled to 91 mesocosms $(1.5 \text{ m}^3 \text{ rock})$ using the same rock samples but with larger size fragments than in the columns. SC declined significantly in both the columns and the mesocosms. Compared to the columns, the peak leachate SC was higher and the temporal decline of SC was not as steady in the mesocosms likely because the mesocosms were in operated in a less controlled environment. In either case, the use of such laboratory and field kinetic tests can involve months to years to obtain results and generally requires kg of rock materials.

 A rapid batch extraction method that can be used to test small quantities of representative materials and that correlates well with field leachates would be of value to identify and manage rock types that release high TDS, and to characterize TDS from different energy extraction activities. For in-field

 determinations, "rapid" could refer to days if samples were shipped to a commercial laboratory, or hours if rock samples could be crushed and sieved, reacted with extractant solutions, and quantified for TDS release based on SC using a portable meter (discussed below). Because ABA parameters such as maximum potential acidity (MPA) and net neutralization potential (NNP) are used by coal companies to characterize overburden and are readily measured, their ability to predict TDS release has been evaluated. Odenheimer et al. (2014) demonstrated that MPA and NNP may be useful to indicate general levels of low, moderate, and high TDS release; however, their semi-quantitative model was based on TDS computed from paste SC for a pulverized rock sample and did not consider upscaled or field-measured leachate characteristics. Modifying a method described by Barnhisel and Harrison (1976) and O'Shay et al. (1990), Orndorff et al. (2010) developed an alternative to the MPA method that used hydrogen peroxide (30% H_2O_2) to oxidize sulfide minerals. They found that the peroxide potential acidity (PPA) was better than MPA as a predictor of TDS release from low-S rocks. However, the potential application of PPA to predict TDS release from a wide range of rock types was not evaluated.

The objectives of this research were to 1) develop and test a batch extraction method to predict TDS release from a range of rock types associated with energy extraction, 2) compare different batch extraction methods to results reported for column leaching tests and field-scale leachate, 3) evaluate those batch extraction methods to determine the most reliable method to quantify TDS release, and 4) identify tracers in leachate that may distinguish coal mining-derived TDS from other energy extraction sources.

- 2. **Materials and Methods**
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2.1 Sample Collection and Preparation

A total of 65 sedimentary rock and coal samples were obtained from multiple sources (Table 1). The samples represent important fossil fuel-bearing strata in the Appalachian Basin, including bituminous coals and gas-producing shales. Eleven rock samples (3 weathered coal refuse, 8 overburden) were provided by Virginia Tech. Equivalent samples had been previously used in laboratory-scale, flow- through unsaturated column leaching experiments (Daniels et al., 2016; Daniels et al., 2014b; Orndorff et al., 2015). Six rock samples (5 overburden, 1 pyrite) were provided by the U.S. Geological Survey

Table 1. Descriptions of the 65 rock samples tested.

 1 2

Table 1. Descriptions of the 65 rock samples tested (continued).

47 48 49

¹ Source: OH = Ohio Geologic Survey; NDA = Penn State Non-disclosure agreement; PA = Pennsylvania Bureau of Topographic and Geologic Survey; USGS = U.S. Geological Survey; $UK = University of Kentucky$; $VT = Virginia Tech$.

U. Refuse = Unleached refuse; W. Refuse = Weathered refuse.

³ WE = Weathering extent: U = Unweathered (partly oxidized while stored indoors but unleached); W = Weathered (partly oxidized and leached outdoors).

⁴ Minerals identified by XRD and are listed in semi-quantitative order of abundance. Ab = albite; Cal = calcite; Chl = chlorite; Dol = dolomite; Gp = gypsum; Gt = goethite; Ill = illite; Jr = jarosite; Kln = Kaolinite; Ms = muscovite; Poi = poitevinite; Py = pyrite; Qtz = quartz; Rz = rozenite; Sd = siderite; Vrm = vermiculite.

CL = Column leachate; FL = Field leachate; FLY = Field lysimeter; FR = Field runoff.

n.a. = not available.

 (USGS). The five overburden samples had been previously characterized and used in laboratory-scale, flow-through column leaching experiments (Hammarstrom et al., 2009; Hornberger and Brady, 2009). The pyrite sample collected from the Bald Eagle Formation during construction of I-99 at the Skytop roadcut in Centre County, PA, along with paired water chemistry measurements had been previously described (Hammarstrom et al., 2005). Four overburden samples were provided by University of Kentucky. Three of these sample materials had been used in field-scale lysimeter studies (mesocosms) (Agouridis et al., 2012; Sena et al., 2014) and all four had also been used in laboratory-scale, flow- through column leaching experiments (Daniels et al., 2016). Twenty unleached coal refuse samples from the roof and floor of the Pittsburgh Coal Formation were collected from drill core materials stored in a repository maintained by the Pennsylvania Bureau of Topographic and Geologic Survey (TopoGeo; Harrisburg, PA). Samples were collected from cores 8009, 8011, 8012, and 8013 that were drilled in Greene County, PA. One unleached Marcellus Shale sample was collected from drill core materials (Sullivan core at 8276 feet) stored by TopoGeo. One unleached Utica/Point Pleasant Shale sample was collected from drill core materials stored in a repository maintained by the Ohio Geologic Survey (Columbus, OH). Eight samples of Utica/Point Pleasant Shale drill cuttings were provided by two gas development companies working in Pennsylvania. Finally, in March 2017, a total of 11 weathered coal refuse samples, 3 coal samples, and 4 coal refuse leachate samples were collected from two coal refuse disposal facilities (referred to as Mine A and Mine B) in western Pennsylvania.

2.2 Rock Type Categorization

48
49 150 51 151 $\frac{53}{1}$ 152 $\frac{55}{56}$ 153 58 154 60 155 Rock samples were sorted into six operational categories: coal $(n=3)$, overburden $(n=17)$, weathered coal refuse (n=14), unleached but oxidized coal refuse (n=20), gas-bearing shale (n=10), and pyrite $(n=1)$ (Table 1). Coal refuse and overburden categories were differentiated based on the definitions in Pennsylvania Code Title 25 (Environmental protection), Chapter 87 (Surface mining coal), Section 87.1 (Definitions) (25 Pa. Code § 87.1) (Commonwealth of Pennsylvania, 2018a). Specifically, overburden is defined as "the strata or material overlying a coal deposit or between coal deposits in its natural state and shall mean material before or after its removal by surface mining". Coal refuse is defined

as "any waste coal, rock, shale, slurry, culm, gob, boney, slate, clay and related materials, associated with or near a coal seam, which are either brought aboveground or otherwise removed from a coal mine in the process of mining coal or which are separated from coal during the cleaning or preparation operations". Shales closer in age and stratigraphic position to coal formations were included in coal refuse or overburden categories. The gas-bearing shale category included only the Utica/Point Pleasant Shale or Marcellus Shale samples. Pyrite included one sample from the Bald Eagle Formation at Skytop roadcut (Hammarstrom et al., 2005).

163 *2.3 Operational Extractions*

Once received, rock samples were freeze-dried using a Labconco FreeZone 4.5 freeze dry system until constant weight was attained. Samples were crushed to < 4.75 -mm using a hydraulic press at 44.5 kN and thereafter with a mortar and pestle until all particles were $\lt 2$ -mm in diameter. Samples were then pulverized using a Spex 8000 ball mill to produce particles < 0.5-mm diameter (passed through No. 35 sieve).

 $\frac{35}{26}$ 170 $\frac{37}{38}$ 171 40 172 42 173 $\frac{44}{15}$ 174 $\frac{46}{47}$ 175 48
49 176 51 $\frac{53}{178}$ $\frac{55}{56}$ 179 57
58 **180** 60 181 Pulverized rock samples were sent to Geochemical Testing, a certified commercial laboratory in Somerset, PA, to conduct three operational extractions and analyze the SC, pH, and solute concentrations of various leachates. A fourth extraction was conducted at Pennsylvania State University to measure strontium isotopes (${}^{87}Sr/{}^{86}Sr$). In Leach 1 (L1), rock samples were reacted with distilled deionized water 173 (DI) under an ambient atmosphere. In Leach 2 (L2), rock samples were reacted with DI water under a 10/90% $CO₂/N₂$ atmosphere. In Leach 3 (L3), rock samples were reacted with 30% H₂O₂ (70% DI) under a 10/90% $CO₂/N₂$ atmosphere. Aside from the differences noted above, the operational procedure for generating leachates followed the same steps. First, 10.00 ± 0.05 g of pulverized rock (<0.5-mm sieve size) was added to an Erlenmeyer 125 mL flask followed by 20 mL of the extractant solution. Addition of DI water in L1 and L2 was done rapidly in one aliquot. Addition of H_2O_2 in L3 was done slowly by adding 1 mL at a time to reduce bubbling caused by oxidation reactions. Flasks were then placed on a shaker table inside a controlled atmosphere apparatus. The lid of the controlled atmosphere apparatus was left open for L1 or sealed for L2 and L3. For the sealed conditions, $10/90\%$ CO₂/N₂ gas was constantly flushed through

the apparatus. All extractions were shaken for 4 hours at 50 rpm and $20-22^{\circ}C$. After 4 hours, each sample was filtered through a 0.45-µm cellulose acetate filter and pH and specific conductance (SC) of the filtrate were measured immediately (Oakton multiparameter PCTestr 35, calibrated with standards and buffers at $20-22^{\circ}$ C). The filtrate was transferred to a 100 mL volumetric flask and DI water was added to dilute the leachate to a final volume of 100 mL for analysis of elemental concentrations. Blank samples were prepared with DI water or H_2O_2 and followed all steps described above.

The strontium leach $(L₄)$ was prepared by extracting the rock samples three times using DI water. First, 2.0 g of pulverized rock (≤ 0.5 -mm) was added to 15 mL of DI water, shaken for 24 hours using a multi-tube vortexer and centrifuged for 20 minutes at 3,000 rpm. The supernantant was removed, 15 mL of DI water was added to the rock pellet, and the extraction was repeated two more times. The three supernatants were combined, filtered (0.45-μm cellulose acetate), and preserved with nitric acid. $87/86$ Sr was separated from leachates using Sr Spec Eichrom Resin and nitric acid (2 N) to yield 0.1 to 1 μ g of 194 strontium. Separated strontium was analyzed on a ThermoFisher scientific Triton Plus thermal ionization mass spectrometer (TIMS) located at Penn State University EESL. Strontium was also extracted from NIST SRM 987 and IAPSO seawater standards and analyzed for ^{87/86}Sr as reference standards.

197 *2.4 Analytical Methods*

A suite of analytes were measured for each of the three leachates (Supporting Information Tables $SI-1 - SI-3$). SC and pH were measured with electrodes submerged in the undiluted leachate. Major 200 elements (Al, Ca, Fe, K, Mg, Mn, Na, P, S, Si) were measured after dilution to 100 mL on a Thermo 201 Scientific iCAP 7400 inductively coupled plasma optical emission spectrometer (ICP-OES). Minor elements (As, Ba, Co, Cu, Li, Mo, Ni, Pb, Se, Sr, Th, Tl, Ti, U, V, Y, Zn, Zr) were measured on an Agilent 7900 inductively coupled plasma mass spectrometer (ICP-MS). Anions (Br, Cl, NO₃, SO₄) were 204 measured on a Dionex DX-120 ion chromatograph (IC). Total inorganic carbon (TIC) was determined by 205 infrared detection after persulfate oxidation (ASTM, 2017). ABA parameters were determined by standard methods (Sobek et al. 1978). Total sulfur was analyzed by dry combustion, and maximum potential acidity (MPA) was calculated by multiplying total S $(\%)$ by 31.25 to obtain g kg⁻¹ CaCO₃

 equivalent. Neutralization potential (NP) was determined by reacting samples with HCl and titrating the effluent with NaOH using methods of Noll et al. (1988), without modification to account for siderite (Skousen et al., 1997). Net neutralization potential (NNP) was calculated by subtracting MPA from NP; negative NNP values imply a potentially acid-producing sample. The above analyses were conducted at Geochemical Testing, Somerset, PA. ${}^{87}Sr/{}^{86}Sr$ ratios were measured on a thermal ionization mass spectrometer (TIMS) located at the Penn State University Energy and Environmental Sustainability Laboratories. Radium isotopes $(^{226}Ra, ^{228}Ra)$ were measured using a small anode germanium detector gamma spectrometer from Canberra Instruments at geometries consistent with internal standards and certified reference materials (UTS-2). After a 21 day equilibration, 226 Ra was calculated from the average activity of Bi-214 (609 keV) and Pb-214 (295 & 351 keV). Direct measurement of ²²⁸Ra were performed using its 228 Ac daughter at 911 keV.

31 220 $\frac{35}{36}$ 222 $\frac{37}{38}$ 223 40 224 42 225 Mineralogy of the rock samples were characterized by X-ray diffraction (XRD) using a PANalytical X'Pert 165 PRO MPD X-ray diffractometer equipped with a PIXcel detector operated in a 1D scanning mode with all channels active. Samples were subjected to Cu K-α radiation from 5 to 70 degree (2θ) at 45 kV and 40 mA. Semi-quantitative analyses were performed using whole pattern fitting in Jade 2010 software from Materials Data Incorporated of Livermore, CA in conjunction with reference files from the International Centre for Diffraction Data PDF4 database. Mineral detection limits were about 3% (m/m) and uncertainty in mineral fractions were \pm 5%.

2.5 Speciation Modeling Methods

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49 228 $\frac{53}{1}$ 230 $\frac{55}{56}$ 231 57
58 **232** 60 233 The PHREEQC 3.0 aqueous speciation model (Parkhurst and Appelo, 2013) was used with input values for leachate data, corrected for dilution (100/volume leachate recovered), to estimate SC by methods reported by McCleskey et al. (2012) and Appelo et al. (2010) as described by Cravotta and Brady (2015). Input data to PHREEQC included the sample temperature, pH, and the mass concentrations (mg/L) of TIC, SO₄, Cl, F, Br, NO₃-N, P, Si, Ca, Mg, Na, K, Li, Fe, Mn, Al, Ba, Sr, and Zn in the filtered 232 leachates. Both methods calculated SC using the same speciated cations and anions $(H^+, Li^+, Na^+, K^+, Cs^+,$ 233 NH_4^+ , Mg^{2+} , Ca^{2+} , Sr^{2+} , Ba^{2+} , F^- , Cl^- , Br^- , SO_4^{2-} , HCO_3^- , CO_3^{2-} , NO_3^- , and OH^-), trace metals $(Al^{3+}$, Fe^{2+} ,

 Fe^{3+} , Mn²⁺, and Zn²⁺), and charged ion pairs (HSO₄⁻, NaSO₄⁻, NaCO₃⁻, and KSO₄⁻), however, the computations used to determine ionic conductivities were different. Briefly, the Appelo et al. (2010) method calculates the ionic conductivity of solute species using ion diffusion coefficients while the McCleskey at al. (2012) method calculates ionic molal conductivities using transport numbers. Both methods sum the ionic conductivity contributions to indicate the solution SC. Additional details on the SC computations are provided in the Supporting Information.

20 241 22 242 24^{24} 243 $\frac{26}{27}$ 244 28
29 245 31 246 33 247 $\frac{35}{36}$ 248 $\frac{37}{38}$ 249 40 250 The concentration of total dissolved solids (TDS) was computed as the sum of the input concentrations of major dissolved constituents (Ca, Mg, Na, K, SO_4 , Cl, CO_3 , NO_3 , SiO_2) (Fishman and Friedman, 1989) plus minor constituents (Sr, Ba, Fe, Al, Mn, Br), assuming that Fe, Al, and Mn formed hydrous oxides (FeOOH, AlOOH, MnOOH) instead of anhydrous compounds. Cravotta and Brady (2015) showed that TDS values computed accordingly were comparable to the laboratory measured residue on evaporation at 180 ºC for mine effluent samples. Osmotic pressure (OP) was computed as the sum of molal concentrations of the same aqueous species used for SC calculations. The OP computation assumes that 1 mol/kg of each ion exerts approximately 1 mOsm/kg osmotic pressure (Haynes et al., 2013). Cravotta and Brady (2015) showed that computed values of OP for mine effluent samples were comparable to standard laboratory measured values of OP using freezing point depression (Kiyosawa, 2003).

2.6 Data for Upscaling Comparisons

Several of the rock samples characterized by operational batch extractions were previously used in laboratory-scale flow-through column experiments or obtained from field sites with paired water samples (Table 1). For upscaling batch extractions to column experiments, sixteen overburden and three weathered coal refuse samples were compared using mass-normalized TDS. As TDS was not reported for 256 column experiments (only SC), a SC-to-TDS conversion factor ($CV = TDS/SC = mg TDS L^{-1}/\mu S \text{ cm}^{-1}$) was calculated for each paired sample using the SC measured in L1 and the corresponding TDS value calculated using the input concentrations for PHREEQC (Supporting Information Table SI-4). The

 cumulative TDS generated in the column experiment following approximately 14 or 40 discontinuous leaching events was calculated according to:

261 Cumulative column leached TDS
$$
\left(\frac{mg TDS}{L}\right) = \frac{\sum_{n=1}^{l} (SC_i \times V_i \times CV)}{V_{total}}
$$
 Eq 1

where, SC_i = SC measured from i-th leach event (μ S cm⁻¹); V_i = volume of each leach event (L); CV = 263 rock-specific SC-to-TDS conversion factor (mg TDS $L^{-1}/\mu S$ cm⁻¹), and V_{total} = total volume of leaches (L).

 Comparisons between field sites and batch experiments were made based on SC, as this parameter was reported for all field samples. A total of ten field sites were included (referred to as Mine A, Mine B, Mines P, KY1, KY2, KY3, KY9, LKFC, BCS3, Skytop), where SC measured from a select number of rock samples were paired with a varied number of SC values measured in the field. For Mine A, 42 records of SC and additional analytes from leachate drains were compared to five weathered coal refuse samples collected from Mine A. For Mine B, 41 records from leachate drains were compared to six weathered coal refuse samples collected from Mine B. For Mines A and B, records were obtained from a field sampling event in March 2017 and from Hydrologic Monitoring Reports (HMRs) submitted by the coal companies to the Pennsylvania Department of Environmental Protection. For Mines P, three records from influent discharges to three Pittsburgh Coal mining/processing plants on active underground mines (Cravotta and Brady, 2015) were compared to 17 unleached coal refuse samples stratigraphically adjacent to the Pittsburgh Coal Formation (TGS1-TGS17, Table 1). Field results for KY1-KY3 are summarized by Sena et al. (2014). For KY1, 199 records from field lysimeters built on top of a valley fill were compared to rock sample KY1 (unweathered overburden). For KY2, 110 records from field lysimeters built on top of a valley fill were compared to rock sample KY2 (unweathered overburden). For KY3, 203 records from field lysimeters built on top of a valley fill were compared to rock sample KY3 (unweathered overburden). For KY9, 18,064 records from a leachate drain at the toe of two valley fills were compared to rock sample KY9 (unweathered overburden). Three records from influent discharges from coal processing plants (Cravotta and Brady, 2015) and 24 records from discharges from abandoned mines in the Lower Kittanning Formation (Cravotta, 2008) were compared to rock sample LKFC (unweathered

 overburden). Six records from influent discharges from active mines (Cravotta and Brady, 2015) and 10 records from discharge samples from abandoned mines in the Lower to Upper Freeport Formations (Cravotta, 2008) were compared to rock sample BCS3 (unweathered overburden). Four records from drainage from the Skytop roadcut collected in May 2004 (Hammarstrom et al., 2005) were compared to rock sample SKYPA (pyrite).

2.7 *Statistical Methods*

Statistical differences between batch extractions $(L1, L2, and L3)$ and upscaled results for 291 selected parameters were evaluated using the Wilcoxon-signed rank test. Outliers were defined as values greater than the 75th percentile plus 1.5 times inter-quartile distance, or values smaller than the $25th$ 293 percentile minus 1.5 times inter-quartile distance. SC and TDS were compared via correlation (Pearson) 294 analyses. Comparisons between TDS from leaches were compared with TDS from column experiments using Pearson correlation and by comparing the fit of our data (R^2) with the line of equality y = x. SC data from the field were compared with SC from batch extractions using an unpaired t-test. Linear regression equations were generated for each rock category and for the full data set. R was used for all statistical analyses (R Core Team, 2016).

299 **3. RESULTS AND DISCUSSION**

300 *3.1 Comparison of Operational Extractions*

Three operational extractions were designed to have varying reactivity with sulfides, carbonates, $siliicates, sulfates, and salts. DI water alone $(L1)$ was presumed to extract weakly-held exchangeable ions,$ salts, hydrolysis products, and high-solubility minerals. The equilibration of the DI extractant with 10% $CO₂$ atmosphere (L2) was hypothesized to promote carbonate dissolution. Although the pH of the L2 extractant blank was less than that of L1, as described below, this level of $CO₂$ did not create significant changes in the chemistries of leachates produced by L1 versus L2. In contrast, the 30% H_2O_2 in L3 307 promoted sulfide oxidation, and the production of sulfuric acid promoted the dissolution of many other minerals. Relationships between SC, TDS, pH, and TIC reflect the major reactions promoted by the extractant solutions. For example, for rocks with high sulfide and low carbonate contents (as determined

by XRD), the increase in SC and TDS after the addition of H_2O_2 (i.e., L1 vs L3) was dominated by 311 production of SO_4^2 and H⁺. As another example, for rocks with low sulfide and high carbonate contents, the increase in SC and TDS after reaction with CO_2 (i.e., L1 vs L2) was controlled by release of Ca^{2+} and 313 $HCO₃$.

314 Chemistry data for all rock samples are provided as Excel files in Supporting Information Tables 315 SI-4 (leachates), SI-5 (blanks), and SI-6 (solids). For the blanks, the median (and range) of pH values were: L1, 6.0 (4.1 – 7.0, n=7); L2, 5.1 (4.4 – 6.3 , n=5); and L3, 2.6 (1.8 – 4.5, n=7). The median measured SC values for the blanks were 13 μ S/cm (5.0 – 85, n=7), 45 μ S/cm (15 – 90, n=5), and 227 μ S/cm (33 – 973, n=7) in L1, L2, and L3, respectively. The median calculated TDS values for the blanks were 26 mg/L (16 – 223, n=7), 30 mg/L (14 – 34, n=5), and 138 mg/L (26 – 339, n=7) in L1, L2, and L3, respectively.

31 322 $\frac{33}{1}$ 323 $\frac{35}{26}$ 324 38 325 40 326 42 327 $\frac{44}{15}$ 328 $\frac{46}{47}$ 329 48
49 330 51 331 $\frac{53}{1}$ 332 $\frac{55}{56}$ 333 58 334 60 335 321 Operational extractions L1, L2, and L3 were compared based on pH, TDS and SC results from all 322 rock samples (Figure 1). The D'Agostino & Pearson normality test showed that the measured values for these parameters did not follow a normal distribution, therefore, comparisons between leachates were made using the Wilcoxon-signed rank test. This test showed that SC (measured and calculated), TDS, and 325 pH values were significantly different (at 95% confidence) in L3 compared with L1 and L2, and that differences between L1 and L2 were not significant (Supporting Information Table SI-7). In general, L3 showed an increase in SC and TDS and a decrease in pH compared with L1 (Figure 1B, D, F). As noted above, this was the result of the oxidation of sulfide minerals promoted by the use of H_2O_2 in L3. The 329 decreased pH promoted the dissolution of carbonate minerals and release of $HCO₃$, $Ca²⁺$, and other ions into solution. Most of the samples showed a substantial increase in SC in L3 compared to L1 (Figure 1B). 331 However, six samples showed only a modest increase in SC (samples touching line of equality in Figure 332 1B). These six samples contained high sulfate and low sulfide contents where the sulfate salts were quickly dissolved by water and the addition of H_2O_2 did not greatly enhance mineral dissolution. Of these six samples, four were unleached coal refuse (TGS 2A, TGS8, TGS10A, and TGD10B), one was weathered coal refuse (TNR2), and one was shale (SOH1). Consistent with statistical paired tests, results

57 337 58 338 $\frac{59}{60}$ 339 **Figure 1.** Comparisons between different batch extractions. Panels A, C and E (n=59) show correlations between Leach 1 (L1) and Leach 2 (L2). Panels B, D, and F (n=63) show correlations between L1 and Leach 3 (L3). Panels G and H (n=63) show correlations between SC and TDS in L1 and L3.

from L1 and L2 were similar (Figure 1A,C,E). However, five samples showed an increase in SC in L2 compared to L1 (Figure 1A). Of these five samples, two were gas-producing shales (SHM2, SHM3), one was unweathered coal refuse (TGS13), and two were sandstone overburden (VA16, WV5). The first three samples had abundant calcite and minor pyrite (Table 1). Although neither mineral was detected by XRD for VA16 or WV5 (Table 1), the two overburden samples had detectible NP and S (Table S3). In any case, the added $CO₂$ in L2 appears to have enhanced carbonate dissolution.

The majority of pH values for L3 were lower than L1 (Figure 1F) because of enhanced sulfide oxidation by $H_2O_2+C_2O_2$. However, a large number of samples (23 of 65) with pH values ranging from 6.5 to 7.9 (samples clustered in upper right of Figure 1F) exhibited little change in pH in L1 versus L3. These samples contained high carbonate and low sulfide contents, with corresponding positive values of NNP 350 (Table 1 and Supporting Information Table S6), and produced enough alkalinity to neutralize the sulfuric acid produced. Of these twenty-three samples, eight were shales, six were unleached coal refuse, two were weathered coal refuse, and seven were overburden.

For the six rock types, median values for pH, SC, TDS, and OP were used to compare L1 and L3 354 (Figure 2). In general, pyrite and unleached coal refuse produced the highest median values for SC and 355 TDS while overburden produced the lowest median values. Median values for these parameters from 356 weathered coal refuse, coal, and shale always ranked in intermediate positions, although the order switched depending on the analyte or extraction method. For instance, the order for TDS (mg/L) in L1 was pyrite $(7,770)$ > unleached coal refuse $(2,430)$ > weathered coal refuse $(1,870)$ > shale $(1,020)$ > coal (375) > overburden (262), while the order in L3 was pyrite $(35,200)$ > unleached coal refuse (8,920) > weathered coal refuse $(6,160) >$ coal $(4,700) >$ shale $(3,360) >$ overburden $(1,080)$. Median values of TDS and SC in L3 were all higher than corresponding proposed regulatory reference levels, 500 mg/L and 300 -500μ S/cm (Cormier et al., 2013a, 2013b; Pond et al., 2008; Timpano et al., 2010), respectively, for all rock types. Except for overburden, median values of OP in L3 were all higher than the regulatory reference level of 50 mOsmol/kg (Commonwealth of Pennsylvania, 2018b).

values are nearly two orders of magnitude greater than the median TDS obtained for gas-bearing shales with our most aggressive extractant (TDS_L3 = 3,360 mg/L). Produced water from hydraulically fractured shale gas wells could encounter extensive small fracture networks equivalent to exceptionally low water-to-rock ratios or could interact with brines that were not within (or preserved) in our samples. This is consistent with other studies that have shown TDS values from batch extractions of Marcellus Shale are much lower than corresponding field produced waters (Chapman et al., 2012; Phan et al., 2015; 382 Rowan et al., 2011; Stewart et al., 2015, 2014; Tasker et al., 2016; Warner et al., 2012).

383 *3.2 Leachate Composition*

The median concentrations of dissolved metals, metalloids, and anions varied based on rock type, weathering extent, and final pH of the extract (Figure 3). Based on sample mineralogy and leachate 386 composition, the principal mechanisms for TDS generation are oxidation of sulfide minerals (with production of H_2SO_4) that promotes increased solubility of metals (e.g., Fe, Al, Mn), dissolution and hydrolysis of carbonate and silicate minerals to neutralize acidity, and dissolution of high-solubility minerals such as sulfates and salts. Because of its high organic carbon content, coal contained relatively low concentrations of metals. Based on XRD (Table 1), coal samples contained quartz, calcite, clays, and pyrite. Concentrations of Si, Ca, Al, Fe, and SO₄ in the coal leachates are consistent with this mineral assemblage.

Weathered coal refuse had been exposed to shallow subsurface weathering for years, while rock 394 cores classified in this study as unleached coal refuse had been exposed to humid air only while archived in core boxes. These differences in weathering extent led to distinct differences in leachate chemistry. Unleached coal refuse released higher alkali metals, notably Na, and higher Cl compared to weathered coal refuse (Figure 3), reflecting that salts had been preserved in storage. Plots of Cl versus Na molar concentrations in both L1 and L3 showed that only the shale samples consistently plotted along the $1:1$ line of equality supporting the assumption of NaCl dissolution (Supporting Information Figure SI-1). Na was also likely sourced from exchange reactions and silicate neutralization. Unleached coal refuse also released higher SO₄, notably in L3, compared to weathered coal refuse, reflecting that some sulfides had

 $\frac{56}{57}$ 405 57 405
58 406 58 406
59 407 60 408 Figure 3. Summary of chemistry for Leach 1 (L1) and Leach 3 for the six rock types. Coal (n=3); Weathered Coal Refuse = W. Ref (n=13 or 14); Unleached Coal Refuse = U. Ref (n=20); Overburden = Overb (n=17); Shale (n=10); Pyrite (n=1). Box plots show median, 25% and 75% quartile ranges. Whiskers show the minimum and maximum values. Outliers (circles) defined as any point at a distance greater than 1.5 times the interquartile range measured from the $75th$ to the $25th$ percentile.

not been oxidized during core storage. Concentrations of transition metals and Se were similar between unleached and weathered coal refuse (Figure 3).

Differences in mineral composition (Table 1) of overburden, shale, and pyrite help explain differences in leachate chemistry. In L3, overburden samples released low alkali metals and chloride, reflecting low entrained salt content, and low SO₄ reflecting low sulfide content. Gas-bearing shale samples released high alkali metals and the highest amounts of Cl and Br, reflecting relatively high salt content, low SO₄ reflecting low sulfide content, and high alkaline earth metals, notably Sr and Ba, reflecting high carbonate content. The sole pyrite sample released the highest amounts of Fe, SO₄, Al, Pb, Zn, and As reflecting high sulfide content.

418 *3.3 Contribution of ionic species to specific conductance*

The relative contributions of ionic species to the SC were calculated for all rock types in L1 and L3 using the method of McCleskey et al. (2012). For L1, the major cationic contributions to SC were Ca^{2+} 421 > Na⁺ > Mg²⁺ > Fe²⁺ > H⁺ > K⁺, and major anionic contributions to SC were SO₄² > Cl⁻ > HCO₃⁻ (Figures 4 and 5). However, the rank of ion contributions to SC differed slightly depending on rock type. For instance, Ca^{2+} was the dominant cation in coal, weathered coal refuse, overburden, and gas-bearing shale, but Na⁺ was the dominant cation in unleached coal refuse and Fe^{2+} was the dominant cation in pyrite. On 425 the other hand, the anionic contributions to SC were dominated by SO_4^2 for all rock types except for gas-426 bearing shale, where Cl was most abundant. The high contribution of Cl \mathbb{R} , Na⁺, and Ca²⁺ to SC in gasbearing shales (Figure 4I, J) provide evidence for salt dissolution. The general contribution of principal cations and anions to SC in L1 were consistent with the mechanisms of TDS generation discussed above. With the addition of H_2O_2 to L3, sulfide oxidation and the consequent release of sulfuric acid 430 became an important mechanism for ion mobilization by mineral dissolution. While the major ions that contribute to SC in L3 were similar to those in L1, the contribution of H^+ increased markedly in L3 (Figure 5). The major cationic contributions to SC in L3 were $H^+ > Ca^{2+} > Fe^{2+} > Na^+ > Mg^{2+} > K^+$. 433 Anionic contributions to SC in L3 were dominated by $SO_4^2 > HSO_4 > Cl^2 > HCO_3$. Na⁺ was an important contributor to SC from unleached coal refuse and gas-bearing shale. The increased release of $Na⁺$ from

 $\frac{51}{52}$ 437 $\frac{52}{53}$ 438 53 438
54 439 55 440 56 441 57 442 **Figure 4.** Ionic contributions to specific conductance in Leach 1 (L1). Median specific conductance (SC), and ionic contributions calculated according to McCleskey et al. (2012). Median pH is shown for each rock category. Left panels show cationic contributions normalized to median SC for each rock category. Right panels show anionic contributions normalized to median SC for each rock category. Coal $(n=3)$; Weathered Coal Refuse = W. Refuse $(n=14)$; Unleached Coal Refuse = U. Refuse $(n=20)$; Overburden $(n=17)$; Shale $(n=10)$; Pyrite $(n=1)$.

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57 449 58 450 Figure 5. Ionic contributions to specific conductance in Leach 3 (L3). Median specific conductance (SC), and ionic contributions calculated according to McCleskey et al. (2012). Median pH is shown for each rock category. Left panels show cationic contributions normalized to median SC for each rock category. Right panels show anionic contributions normalized to median SC for each rock category. Coal $(n=3)$; Weathered Coal Refuse = W. Refuse $(n=14)$; Unleached Coal Refuse = U. Refuse $(n=20)$; Overburden $(n=17)$; Shale $(n=10)$; Pyrite $(n=1)$.

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gas-bearing shale with L3 compared to L1 is consistent with silicate mineral decomposition combined with salt dissolution. Cl was an important contributor to SC of L3 only from gas-bearing shale, where the 453 SO₄² release was greater than Cl. Mg²⁺ and HCO₃ were important contributors to SC only from overburden. Although pH of L3 remained near-neutral for the gas-bearing shale and overburden, the 455 increased release of SO_4^2 , Ca^{2+} , and Na⁺ with L3 compared to L1 demonstrates the importance of mineral decomposition in conjunction with pyrite oxidation, acidification, and neutralization.

457 *3.4 Upscaling from lab to field*

An important part of this study was to test the capability of the proposed rapid batch extractions on 10 g samples for predicting TDS release from coal refuse and overburden of larger size and at longer time scales. Available information from previous column studies and water quality data from ten field sites were compared with our batch extractions. Seventeen overburden samples and three weathered coal 462 refuse samples (Table 1) were previously analyzed in column studies(Agouridis et al., 2012; Daniels et 463 al., 2009; Daniels et al., 2016; Daniels et al., 2014b; Hornberger and Brady, 2009; Odenheimer et al., 2015; Sena et al., 2014). In general, all column studies maintained unsaturated conditions with simulated rainfall (pH 4.6) leaching events. The number and frequency of leach events, the rainfall volume, and the rock mass in the columns varied between experiments conducted by Daniels et al. (2016) versus Hornberger et al. (2009). Daniels et al. (2016) constructed columns with 1,200 cm³ (\sim 1,800 g) of rock and applied 125 mL of synthetic rain water twice a week for a total of 40 leach events. Hornberger et al. (2009) constructed columns with 1,300 to 2,100 grams of rock and applied 190 to 650 mL of synthetic rain water once a week for a total of 14 leach events. As described below, the overall water volume (sum 471 of individual leaches) to rock mass ratio strongly controls leachate chemistry. As these two groups of 472 researchers ultimately used similar water-to-rock ratios in their column studies, results from all studies are comparable when normalized to rock mass. In general, SC in the column leachates started at the highest 474 values, declined in an exponential manner, and then approached an asymptotic minimum. Comparisons with batch experiments were made based on cumulative TDS calculated from the column experiments. 476 Column leachate data were compiled as SC and then converted to TDS, based on rock-specific

 Figure 6. Cumulative total dissolved solids (TDS) calculated for column experiments (calculated using Eq. 1) versus TDS measured in Leach 1 (L1) and Leach 3 (L3). For column experiments with replicates, symbols represent mean values and error bars represent standard deviation. Error bars smaller than the symbol size are not shown. Weathered coal refuse $= W$. Refuse (n=3). Overburden samples donated by $USGS = Overburden (USGS)$ (n=5), overburden samples from other sources = Overburden (Other) $(n=11)$.

 TDS from batch extraction L3 tended to overpredict cumulative TDS calculated from the column experiments (Figure 6B; Supporting Information Table SI-8). These results suggest that even multiple column leaching events cannot achieve the extractive strength of $H_2O_2+CO_2$ used in batch extraction L3. Furthermore, these results suggest that multiple discontinuous rainfall leaching events do not substantially

increase the extractive strength of synthetic rain (or physical access to additional reactive sites), and that the cumulative water-to-rock ratio exerts greater control on leachate chemistry for water extractions.

In contrast to the column experiments, field SC was better correlated to SC from batch extraction L3 versus L1 (Figure 7; Supporting Information Table SI-9). Field data and batch extractions were 503 compared for nine of the 10 field sites using an unpaired t-test (not enough data were available to test the 504 other three sites). Seven of the nine sites showed no statistical differences between field SC and L3 SC. We note that field SC values from rock disposal facilities (a.k.a. excess spoil fills pursuant to SMCRA) 506 change over time (Evans et al., 2014), and that the 'age' of the rock/leachate could not be controlled in this study. However, we chose to analyze all these sites together because they represent the largest, most 508 analytically consistent data set available for addressing our research objectives. In contrast, five of the 509 nine sites showed significant statistical differences between field SC and L1 SC, where SC results from L1 underestimated the field SC.

511 Collectively these upscaling comparisons suggest that weathering in the field is influenced by acid-formation and neutralization reactions that produce greater solute concentrations than simple 513 dissolution of soluble salts and exchangeable ions by water alone. Column leaching experiments produce high SC in the first leaches but values decline rather quickly. Scaling up from column experiments to field sites is challenging for a number of reasons. One obvious issue is that coal refuse disposal fills often 516 contain millions of cubic meters of rock such that the rock-to-water ratio is dramatically greater in the field as compared to tens of pore volumes eluted through laboratory columns. Unlike column experiments, water percolating through rocks stored in disposal fills may encounter multiple and much longer flow paths such that the water encounters more 'fresh' reactive material. Water may migrate through these rocks much slower and encounter many more wetting-and-drying cycles as compared to 521 column experiments such that the field leachates oxidize more sulfides, generate a lower pH, and solubilize more metals. Rocks in disposal fills may also disaggregate over long periods of time, effectively increasing the rock-to-water ratio.

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27 527 28 528 $\frac{29}{20}$ 529 $\frac{30}{31}$ 530 $\frac{31}{32}$ 531 33 532 533 **Figure 7.** Relationships between specific conductance (SC) measured in the field and in A) Leach 1 (L1) and B) Leach 3 (L3). Symbols represent mean values and error bars represent standard deviation. For SC measured in the field: Mine A (n=42); Mine B (n=41); Mines P (n=3); KY1 (n=199), KY2 (n=110), KY3 $(n=206)$, KY9 (n=18,064); LKFC (n=25); BSC3 (n=16); Skytop (n=4). For SC measured in L1: Mine A (n=5); Mine B (n=6); Mines P (n=17); KY1 (n=3); KY2 (n=3); KY3 (n=3); KY9 (n=1); LKFC (n=3); BSC3 (n=3); Skytop (n=8). For SC measured in L3: Mine A (n=5); Mine B (n=6); Mines P (n=17); KY1 $(n=3)$; KY2 $(n=3)$; KY3 $(n=3)$; KY9 $(n=1)$; LKFC $(n=2)$; BSC3 $(n=2)$; Skytop $(n=14)$.

 $\frac{38}{10}$ 535

3.5 Correlations between SC and ABA parameters

 536 $\frac{42}{42}$ 537 $\frac{44}{45}$ 538 47 539 49 540 $\frac{51}{10}$ 541 $\frac{53}{54}$ 542 56 543 58 544 Acid-base accounting (ABA) parameters are used to identify and segregate rocks with high acid generation potential (or blend with rocks with high alkalinity). The use of ABA parameters to identify TDS release potential would be cost-effective for coal mine operators, provided that the TDS predictions based on ABA were accurate. Of all the correlations between ABA parameters and analytes measured in L1 and L3 extracts, maximum potential acidity (MPA) displayed the most promising correlations with SC from L3 (Table 2). It was anticipated that $MPA + NP$ might better predict TDS release, but this did not produce an improved relationship. While MPA is certainly correlated with TDS release, correlation coefficients were not high for all rock types and notably low for weathered coal refuse, plus different linear regression coefficients (slope and intercept) were indicated for different rock types (Supporting

 Information Figure SI-2). Therefore, an additional measure such as batch extraction L3 to measure TDS release potential would still be recommended.

Table 2. Correlations between TDS and ABA parameters. MPA=Maximum Potential Acidity;

 $SC = Specific conductance; r = correlation coefficient.$

Correlations between XRD-based mineral contents, ABA parameters, and selected leachate

chemistry parameters (Supporting Information Table SI-11) confirmed that samples containing sulfide

and sulfate minerals had higher total S, and samples containing calcite and dolomite had higher NP. The strongest predictor of leachate salinity (SC, TDS, or OP) was the total S content and the presence of sulfur 555 minerals. These correlations support the hypothesis that Leach 1 liberates sulfur and iron mainly from 556 iron sulfate minerals formed by prior oxidation of pyrite. Identification of iron sulfide minerals does not seem to be particularly informative for predicting water chemistry. Generally, the significance of 558 correlations between ABA parameters and salinity parameters increases for Leach 3 (which seems to mobilize Ca from carbonates, whereas Leach 1 mobilizes Ca from gypsum).

560 *3.6 Rock Type Fingerprinting*

In certain situations, the ability to distinguish the source of TDS contamination is valuable. This 562 becomes more challenging in the Appalachian Basin where coal mining, conventional oil and gas (O&G) 563 production, and unconventional gas production all coexist. The ability to distinguish TDS inputs from 564 coal versus O&G activities is important for establishing corrective and preventive actions. Bromide, strontium isotopes (${}^{87}Sr/{}^{86}Sr$), and radium isotopes (${}^{228}Ra/{}^{226}Ra$) have all been used to identify the addition 566 of O&G produced water into freshwater systems (Chapman et al., 2012, Rowan et al., 2011, 2015; Jonson 567 et al., 2015, Warner et al., 2012). As noted above, shales produced Na-Cl waters that were generally distinct from Ca-SO₄ waters produced from coal-associated rocks. Therefore, these potential geochemical tracers combined with chloride were examined for fingerprinting purposes.

 $\frac{44}{15}$ 571 $\begin{array}{c}\n46 \\
47 \quad 572\n\end{array}$ 48
49 573 51 574 $\frac{53}{11}$ 575 $\frac{55}{56}$ 576 58 577 60 578 Leachate chemistry from gas-producing shales (nine samples from Utica/Point Pleasant Shale 571 plus one sample from Marcellus Shale) compared to coal mining-associated materials showed that differentiation with Cl versus C/SO_4 molar ratio and Cl versus ${}^{87}Sr/{}^{86}Sr$ isotope ratio are the most effective tools for source identification (Figure 8). Br was not a robust tracer in this study because it was below detection in most samples from coal-bearing strata (e.g., 24 of 65 samples had measurable Br in 575 L1; 9 of 65 samples had measurable Br in L3). Where Br values were above detection limits, Cl/Br and Cl 576 were greater in the gas-bearing shale samples compared to the coal-associated rocks. Ra isotopes were not 577 effective tracers because of relatively large and overlapping variances in both total Ra activity and the ²²⁸Ra^{/226}Ra isotope ratio for each rock type (Supporting Information Figure SI-3). ⁸⁷Sr^{/86}Sr isotope ratios

al., 2016; Chapman et al., 2012; Capo et al., 2014; Blondes et al, 2017). The median $87Sr/86Sr$ isotope ratio for the Utica/Point Pleasant Shale from 26 samples is 0.7110 while the $25th$ to $75th$ percentiles range from 0.7109 to 0.7114 (Tasker et al., 2019). It must be noted that insufficient amounts of the shale samples

used in this study were available for conducting ${}^{87}Sr/{}^{86}Sr$ isotope measurements. Instead, all gas-

from over 133 samples is 0.7112 while the $25th$ to 75th percentiles range from 0.7110 to 0.7114 (Phan et

B)

D)

 $10¹$

 10^{-1}

 10^{-2}

 $\overline{\bigcirc}$

Br/Cl L3 (mol/mol)

 $10⁰$

CI_L1 (mmol/L)
producing shale samples in Figure 8C,D are represented with the shaded region showing these $25th$ to $75th$ percentiles range. Using these well-constrained values for ${}^{87}Sr/{}^{86}Sr$ isotope ratios combined with Sr/Ca 596 molar ratios or Cl concentrations, differentiation of TDS from gas-producing black shales and coalassociated rocks is possible.

598 **4. CONCLUSIONS**

Increased salinization of fresh water resources is a growing concern even in water-rich regions such as the Appalachian Basin. Management of activities and industries that release TDS could reduce this problem. For coal mining, segregation and isolation of rocks that produce high levels of TDS is one 602 obvious management strategy. To implement this strategy, a rapid and simple method to identify these rocks by quantifying TDS release is required. In regions with coal mining and other sources of TDS (e.g., 604 coal-bed methane, oil & gas development, road brining), source identification could also help reduce TDS release and enhance the information available to decision makers.

 $\frac{46}{47}$ 613 In this study, the mass of TDS released from sedimentary rocks (65 samples) was measured in laboratory batch extractions and compared to upscaled results from flow-through columns (19 samples) and field measurements (35 samples paired to 10 sites each with multiple field records). The pH of extractant blanks used for the batch tests decreased in the order DI (6.0), $DI+CO₂ (5.1)$, and $H₂O₂+CO₂$ (2.6), which indicated the 30% H_2O_2 was mildly acidic as well as an oxidant. The DI extractant was effective for mobilizing soluble SO_4 and Cl salts, which are predominant sources of TDS upon initial wetting of crushed rock. The DI+CO₂ extractant increased the weathering of carbonates present in some samples, but did not significantly increase TDS production compared to the DI extraction when considering the whole set of samples. The $H_2O_2+CO_2$ extractant increased the weathering of sulfides (and 615 carbonates) and resulted in the greatest TDS production and lowest pH values. When accounting for the mass of rock-to-volume of extractant, TDS measured in batch extractions was strongly correlated to 617 cumulative TDS calculated from column experiments. TDS measured in batch extractions using 30% $H₂O₂$ under 10% CO₂ was higher and poorly correlated to cumulative TDS calculated from upscaled

column experiments. Results suggest that the cumulative water-to-rock ratio controls leachate chemistry in batch extractions using DI water or flow-through configurations using synthetic rain.

 Because all ions were not measured in field samples such that TDS concentrations could not be calculated, batch extractions and field measurements were compared based on SC. In contrast to column experiments, field SC was better correlated to SC measured from $H_2O_2+CO_2$ extractions versus DI extractions. The field SC and SC from $H_2O_2+CO_2$ extractions were statistically indistinguishable for 7 of 9 paired data sets while SC from DI extractions underestimated field SC in 5 of 9 cases. Compared to column leaching over months or waiting until mined rock begins weathering in the field, the batch extractions of small samples are efficient and informative. The small sample size used in batch tests permits testing of specific lithologies or strata. Results were comparable among the rapid batch tests and longer-term laboratory or field data sets. Upscaling comparisons suggest that (1) weathering reactions in the field are more aggressive than DI water or synthetic rainwater extractants used in batch or column tests, and (2) a batch extraction method utilizing 30% H_2O_2 (which is mildly acidic without CO_2 enrichment) could be effective for identifying rocks that will release high amounts of TDS.

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 $\frac{17}{18}$ 639 20 640 22 641 24 642 $\frac{26}{27}$ 643 28
29 644 31 645 $\frac{33}{1}$ 646 **Author Contributions:** Luis Castillo Meza – formal analysis, writing – original draft, writing – review and editing; Charles Cravotta – formal analysis, writing – original draft, writing – review and editing; Travis L. Tasker – formal analysis; Nathaniel Warner – writing – original draft, writing – review and editing; Lee Daniels – formal analysis, writing – review and editing; Zenah Orndorf – formal analysis, writing – review and editing; Tim Bergstresser – formal analysis; Amy Douglass – formal analysis; George Kimble – formal analysis; Joelle Streczywilk – formal analysis; Chris Barton – formal analysis, writing – review and editing; Stephanie Fulton – formal analysis, writing – review and editing; Aaron Thompson – formal analysis, writing – review and editing; and William D. Burgos – supervision, writing $-$ original draft, writing – review and editing.

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SUPPORTING INFORMATION

Batch Extraction Method to Estimate Total Dissolved Solids (TDS) Release from Coal Refuse and Overburden

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Eleven tables, three figures, and expanded version of materials and methods:

Table SI-1. Summary of operational batch extractions and associated measurements. (this file)

Table SI-2. Summary of analytical methods used for leachates. (this file)

Table SI-3. Summary of analytical methods used for solids. (this file)

Table SI-4. Sample descriptions and associated PHREEQC input and output data for rapid leach samples: type 1 (deionized water), type 2 (10% CO2), and type 3 (30% H2O2+10% CO2). (Excel file)

Table SI-5. Sample descriptions and associated PHREEQC input and output data for blank samples: type 1 (deionized water), type 2 (10% CO2), and type 3 (30% H2O2+10% CO2). (Excel file)

Table SI-6. Chemical composition of rock samples used for rapid leach tests. (Excel file)

Table SI-7. Summary of statistical comparisons (Wilcoxon signed-rank test) between selected properties (measured or calculated) in the three batch extractions. SC_L - Specific conductance measured in leachates; SC_CM – Specific conductance calculated by McCleskey method; TDS – Total dissolved solids (calculated); Sig. Diff. – Significantly different in a 95% confidence interval. (this file)

Table SI-8. Correlations between TDS from column experiments and TDS from batch extractions. TDS=Total dissolved solids; $r =$ correlation coefficient; $R^2 =$ regression coefficient. (this file)

Table SI-9. Summary of statistical comparisons (Unpaired t-test) between specific conductance measured in batch extractions and corresponding field sites. SC_L1 - Specific conductance measured in Leach 1; SC_L3 - Specific conductance measured in Leach 3; Sig. Diff. – Significantly different at a 95% confidence interval. (this file)

Table SI-10. Summary of field chemistry and leachate results for the 10 paired field sites. (Excel file)

Table SI-11. Spearman rank correlation coefficient (r) matrix for XRD, acid-base account parameters, and leachate chemistry. (Excel file)

Figure SI-1. Plots of Cl versus Na molar concentrations in both L1 and L3 as compared to 1:1 line of equality. (this file)

Figure SI-2. Relationships between specific conductance (SC) in leach 3 (H2O2+CO2) and acid base accounting (ABA) parameters. (this file)

Figure SI-3. Total Ra and ²²⁸Ra/²²⁶Ra isotope ratios for the six rock types. Coal (n=3); Weathered Coal Refuse = W. Ref (n=14); Unleached Coal Refuse = U. Ref (n=20); Overburden = Overb (n=17); Shale $(n=10)$; Pyrite $(n=1)$. Box plots show median, 25% and 75% quartile ranges. (this file)

Supporting Information-1 – Expanded version of Materials and Methods

Sample Preparation

Samples were freeze-dried using a Labconco FreeZone 4.5 freeze dry system until constant weight was attained (~24 h). Samples were crushed to >4.75 mm using a hydraulic press at 44.5 kN of force and thereafter with a mortar and pestle until all particles were < 2 mm in diameter. Samples were further pulverized using a Spex 8000 ball mill to produce particles < 0.5 mm diameter (passed through No. 35 sieve).

Water Samples

Six water samples were collected from leachate drains of refuse piles at Mines A and B in March 2017. Conductivity and pH were measured in the field using a HACH HQ40 portable multimeter. Historical water quality data from these same leachate drains were compiled from Hydrologic Monitoring Reports (HMRs) submitted by the coal companies to the Pennsylvania Department of Environmental Protection (PA DEP). Historical water quality data from these same leachate drains were compiled from Hydrologic Monitoring Reports (HMRs) submitted by the coal companies to the Pennsylvania Department of Environmental Protection (PA DEP). HMR data spanned from 10/28/17 to 03/30/18 for Mine A, and from 10/07/14 to 10/26/17 for Mine B. Water samples were collected from the Skytop roadcut on May 2004.

Operational Extractions

Pulverized rock samples were sent to Geochemical Testing, a certified commercial laboratory in Somerset, PA, to conduct four operational extractions and analyze the various leachates. A fifth extraction was conducted at Pennsylvania State University to measure strontium isotopes (${}^{87}Sr/{}^{86}Sr$). Extractions are summarized in Table SI-1.

Table SI-1. Summary of operational batch extractions and associated measurements.

In Leach 1, rock samples were reacted with distilled deionized water (DI) under an ambient atmosphere. In Leach 2, rock samples were reacted with DI water under a $10/90\%$ CO₂/N₂ atmosphere. In Leach 3, rock samples were reacted with a 30% H_2O_2 (70% DI) under a 10/90% CO_2/N_2 atmosphere. Aside from the differences noted above, the operational procedure for generating the three leaches followed the same steps. First, $10.00\pm0.05g$ of pulverized rock (<0.5-mm sieve size) was added to a 125 mL Erlenmeyer flask followed by 20.0 mL of the extraction solution. Addition of DI water in Leach 1 and Leach 2 was done rapidly in one aliquot. Addition of the H_2O_2 solution in Leach 3 was done slowly by adding 1.0 mL at a time to reduce effervescence of the reaction and loss of sample. The Erlenmeyer flasks were then placed on a shaker table inside a controlled atmosphere apparatus. The lid of the controlled atmosphere apparatus was left open for Leach 1 or sealed for Leach 2 and Leach 3. For the sealed conditions, 10% CO₂ and 90% N₂ gas was constantly flushed through the apparatus. All extractions were shaken for 4 hours at 50 rpm at room temperature. After the 4 hour reaction period, each sample was filtered through a 0.45-µm cellulose acetate filter and pH and specific conductance (SC) of the filtrate were measured immediately. The filtrate was transferred to a 100 mL volumetric flask and DI water was added to a final volume of 100 mL. This diluted sample was distributed into different containers for

further analysis. Multiple blank samples were prepared with DI water or H_2O_2 and followed all steps described above.

Triple acid digestions were accomplished according to the standard method ASTM D 6357-11. Briefly, 2.5 g of pulverized rock (<0.5-mm) were ashed using a Thermolyne FA1740 furnace at 500 ^oC. Thereafter 0.5 g of ash was mixed with 20 mL of aqua-regia and 20 mL hydrofluoric acid in a beaker. The mixture was then heated to dryness followed by addition of 1 mL of concentrated nitric acid and 20 mL of DI water. After heating for 1 h at 100 $^{\circ}$ C and cooling to room temperature, the solution was diluted to 100 mL using DI water. The solution was then analyzed by inductively couple plasma mass spectrometer (ICP-MS) or inductively coupled plasma atomic emission spectroscopy (ICP-OES) following EPA methods 6010 and 6020 respectively.

Strontium leach (L4) was prepared by extracting the rock samples three times using DI water. First, 2.0 g of pulverized rock (<0.5-mm) and 15 mL of DI water were added into a 50 mL metal free plastic tube and shaken for 24 hours using a VWR Multi-tube Vortexer. Tubes were spun for 20 minutes at 3000 rpm using an Eppendorf 5810 R centrifuge and the supernatants were transferred to a new 50 mL tube. For the second extraction, 15 mL of DI water was added to the solid pellet, and the same steps described for the first extraction were followed. The same procedure was repeated for the third extraction with the exception that the tubes were shaken for 12 h. The three supernatants were combined and filtered (0.45-µm cellulose filter), preserved with nitric acid and kept at 4° C until analyzed.

Analytical Methods

Table SI-2. Summary of analytical methods used for leachates.

¹ IC=Ion chromatography; ICP-MS=Inductively couple plasma mass spectrometer; ICP-OES=Inductively coupled plasma atomic emission spectroscopy; TIMS=Thermal ionization mass spectrometry

Extracts from L1, L2, and L3 were filtered, and SC and pH were measured in the filtrate using an Oakton multiparameter PCTestr 35. Extracts were then diluted to 100 mL with DI water and the volume was split for the further analyses. TIC was measured following the ASTM D4839 method using an O.I Analytical 1010 TOC analyzer attached to an O.I. Analytical model 1051 autosampler. Hardness was calculated using mass concentration values of Ca and Mg (measured by ICP-OES).

Major and minor elements were meassured on a Thermo Scientific iCAP 7400 inductively coupled plasma optical emission spectrometer (ICP-OES), Agilent 7900 inductively coupled plasma mass spectrometer (ICP-MS), and Dionex DX-120 ion chromatography (IC) with AS22 and AG22 separation and guard columns (4 mm) at Geochemical Testing Lab (Somerset, PA).

Strontium from an aliquot of leach 4 (L4) containing approximately 100 to 800 ng of strontium was separated using Eichrom resin. Yield checks confirmed greater than 98% strontium recovery. 87 Sr/ 86 Sr was measured on a thermal ionization mass spectrometer (TIMS) located at the Penn State University Energy and Environmental Sustainability Laboratories. NIST SRM 987 and IAPSO seawater standards were also separated for strontium and analyzed concurrently with the samples to ensure data quality. The precision of the NIST standard during analysis was 0.7102599 ± 0.000009 (2 x Standard Error). To address any mass interference from ⁸⁷Rb after strontium separation, samples loaded onto filaments for the TIMS were heated past the ionization temperature of the Rb but below the ionization temperature of Sr. This removes some of the residual Rb in the samples. If 85 Rb is detected above its background concentration in the samples during analysis, the ⁸⁷Rb is estimated based on the natural abundance of ⁸⁵Rb and ⁸⁷Rb, and the ⁸⁷Sr/⁸⁶Sr ratio is calculated (⁸⁷Sr = ⁸⁷Total – ⁸⁷Rb). This correction was applied to 7 of the 23 analyzed samples.

Neutralization potential (NP) was determined as the amount of acid neutralized by the sample in CaCO₃ equivalents expressed as g/kg (Mg /1000 Mg of rock). Total sulfur (%) was analyzed using a LECO 628 analyzer equipped with sulfur add-on module, following the directions of the ASTM D423917 method. Maximum potential acidity (MPA) was calculated from the sulfur content assuming complete oxidation of pyrite and neutralization of all generated acidity, according to the following reaction²:

$$
FeS_2 + 2CaCO_3 + 3.75O_2 + 1.5H_2O \Leftrightarrow Fe(OH)_3 + 2SO_4^{-2} + 2Ca^{+2} + 2CO_2
$$
 Eq SI-1

Therefore, after stoichiometry equivalences, MPA was calculated as:

$$
MPA \left(g \frac{cac_{3}}{kg}\right) = S(96) * 31.25
$$
 Eq SI-2

where *S* is the total sulfur concentration (weight percent), and 31.25 is a stoichiometric conversion factor based on Eq SI-1. Net neutralization potential (NNP) in units of (g CaCO \sqrt{kg}) was calculated by subtracting MPA from NP. The MPA computation (and that for NNP) assumes the acidity produced from 1 mol FeS₂ (64 g of S) is neutralized by 2 mol CaCO₃ (200 g) (Cravotta et al., 1990). On this basis, 31.25 g of CaCO₃ will neutralize the acidity from 1,000 g of rock that contains 1.0 weight percent (%) pyritic sulfur.

Radium isotopes (226 Ra, 228 Ra) were measured using a small anode germanium detector gamma spectrometer from Canberra Instruments at geometries consistent with internal standards and certified reference materials (UTS-2). After a 21 day equilibration, 226 Ra was calculated from the average activity of Bi-214 (609 keV) and Pb-214 (295 & 351 keV). Direct measurement of ²²⁸Ra was performed using its 228 Ac daughter at 911.16 keV.

Sediment mineralogy was characterized by qualitative X-ray diffraction (XRD) using a PANanalytical X'Pert 165 PRO MPD located in the Materials Characterization Lab at Pennsylvania State University. The X-ray diffractometer ran from 5-70 degrees 2-theta at a power setting of 45 kV and 40 mA, with a PIXcel detector that was operated in line scanning mode with an active length of 3.34 degrees. Incident side set-up consisted of a 1/4 degree divergence slit and a 1/2 degree anti-scatter slit, and 0.04 radians Soller slits. The diffracted side utilized a 1/4 degree receiving slit, 0.04 radian Soller slits, and a Ni filter. The collected data were then analyzed using JADE for phase identification.

Speciation Modeling Methods

The PHREEQC 3.0 aqueous speciation model (Parkhurst and Appelo, 2013) was used with input values for effluent data to estimate SC by methods reported by Appelo et al. (2013) and McCleskey et al. (2012). Input data to PHREEQC included the sample temperature, pH, and the mass concentrations (mg/L) of total inorganic carbon (TIC), SO₄, Cl, F, Br, NO₃-N, P, Si, Ca, Mg, Na, K, Li, Fe, Mn, Al, Ba, Sr, and Zn in the effluent after filtration $(< 0.45 \mu m$ pore size). Both methods calculate SC using the same speciated cations and anions $(H^+, Li^+, Na^+, K^+, Cs^+, NH_4^+, Mg^{2+}, Ca^{2+}, Sr^{2+}, Ba^{2+}, F^-, Cl^-, Br^-, SO_4^{2-},$ HCO_3^- , $CO_3^2^-$, NO_3^- , and OH $^-$), trace metals $(Al^{3+}$, Fe^{2+} , Fe^{3+} , Mn^{2+} , and Zn^{2+}), and charged ion pairs $(HSO_4^-$, Na SO_4^- , Na CO_3^- , and KSO_4^-). However, the computations used to determine ionic conductivities are different.

Briefly, the Appelo et al. (2013) method calculates the ionic conductivity $(\lambda_{0,i})$ of the above solute species using diffusion coefficient (Dw), ionic charge (z), Faraday's constant (F), gas constant (R), and absolute temperature (T) (equation SI-3)

$$
\lambda_{0.i} = \frac{z^2 F^2}{RT} D_w
$$
 Eq SI-3

Then, conductance (κ) at sample temperature is calculated as the sum of the individual ionic conductivities multiplied by the speciated concentration (m) and the activity coefficient (γ_{sc}) using equation SI-4:

$$
\kappa = \sum (\lambda_{0,i} \gamma_{sc} m_i) \tag{Eq S1-4}
$$

The McCleskey at al. (2012) method calculates ionic molal conductivities (λ_i) at sample temperature (T) and ionic strength (I) using equation SI-5.

$$
\lambda_i = \lambda^0(T) - \frac{A(T)I^{0.5}}{1 + BI^{0.5}}
$$
 Eq SI-5

where I is calculated as equation SI-6

$$
I = 0.5 \sum m_i z_i^{0.5}
$$
 Eq SI-6

For the above equations, *i* is the ion and *z* is its charge, and λ^0 and *A* are functions of temperature and *B* is an empirical constant. McCleskey et al. (2012) then compute the conductance at sample temperature as the sum of the products of molal ionic conductivity (λ_i) and the molal concentration of each of the species (*mi*) (equation SI-7)

$$
\kappa = \sum \lambda_i m_i \qquad \qquad \text{Eq SI-7}
$$

Individual contributions to the SC, expressed as the transport number (*t*), were calculated from the conductivity determined by McCleskey at al. (2012) method using equation SI-8

$$
t_i = \frac{\lambda_i m_i}{\kappa} \tag{Eq S1-8}
$$

Both of the above-cited methods indicate the computed conductance at the sample temperature. Because the temperature in the laboratory was 25 °C, the computed conductance for each sample equals its specific conductance at 25 ºC (SC). For conductivity estimates at other temperatures, the SC can be calculated assuming the conductivity changes approximately 2.1 percent per degree C as reported by McCleskey at al. (2012) (equation SI-9):

$$
SC = \frac{\kappa(T)}{1 + 0.021 (T - 25)} \tag{Eq SI-9}
$$

The data used as input to, or computed as output from, PHREEQC 3.0 were also used to compute total dissolved solids (TDS) and osmotic pressure (OP). The TDS was computed as the sum of the input concentrations of major dissolved constituents (Ca, Mg, Na, K, SO₄, Cl, CO₃, NO₃, SiO₂) (Fishman and Friedman, 1989, p. 437-438) plus minor constituents (Sr, Ba, Fe, Al, Mn, Br), in mg/L, assuming that Fe, Al, and Mn formed hydrous oxides (FeOOH, AlOOH, MnOOH) instead of anhydrous compounds. Cravotta and Brady (2015) showed that TDS computed accordingly was comparable to the laboratory measured residue on evaporation at 180 ºC for mine effluent samples.

Osmotic pressure (OP) was computed as the sum of molal concentrations of the same aqueous species used for conductivity calculation. The OP computation assumes that 1 mol/kg of each ion exerts approximately 1 mOsm/kg osmotic pressure (Haynes et al., 2013). Cravotta and Brady (2015) showed that the OP computed accordingly was comparable to the laboratory measured OP for mine effluent

samples. Measured OP normally is determined using freezing point depression, by which an Osmol is defined as the number of moles of a solute required to lower the freezing point of 1 kg of water by 1.858⁰C (Kiyosawa, 2003).

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Statistical Methods

One of the goals of our study was to compare the different leaching methods. For this evaluation, we first identified outliers by comparing results of SC measured in the extraction leaches with their corresponding SC calculated by McCleskey method¹⁹ (results not shown). Thereafter, statistical differences between leachates for selected parameters (SC, TDS, and pH) were evaluated using the Wilcoxon signed-rank test. SC and TDS were compared via correlation (Pearson) analyses. In addition, linear regression equations were generated for each rock category and for the full data set. The next goal of our study was to compare our batch extractions with column studies and field data. Comparisons between TDS from leaches against TDS data derived from column experiments were made using the Pearson correlation and by comparing the fit of our data (R^2) with line Y = X. Finally, SC data from the field was compared with SC results of our leaches using antwo sample t-test.

Table SI-7. Summary of statistical comparisons (Wilcoxon signed-rank test) between selected properties (measured or calculated) in the three batch extractions. SC_L - Specific conductance measured in leachates; SC_CM – Specific conductance calculated by McCleskey method; TDS – Total dissolved solids (calculated); Sig. Diff. – Significantly different in a 95% confidence interval.

Comparison	SC L		SC CM			TDS	pH	
	P value	Sig. Diff.		P value Sig. Diff. P value Sig. Diff.			P value	Sig. Diff.
$L1$ vs $L2$	0.3477	N	0.7987	N	0.0715	N	0.8077	
$L1$ vs $L3$	< 0.0001		< 0.0001		< 0.0001		< 0.0001	
$L2$ vs $L3$	< 0.0001		< 0.0001		< 0.0001		< 0.0001	

Table SI-8. Correlations between TDS from column experiments and TDS from batch extractions. TDS=Total dissolved solids; $r =$ correlation coefficient; $R^2 =$ regression coefficient.

Table SI-9. Summary of statistical comparisons (Unpaired t-test) between specific conductance measured in batch extractions and corresponding field sites. SC_L1 - Specific conductance measured in Leach 1; SC_L3 - Specific conductance measured in Leach 3; Sig. Diff. – Significantly different at a 95% confidence interval.

n.d. – not determined because of insufficient data

Figure SI-1. Plots of Cl versus Na molar concentrations in both L1 and L3 as compared to 1:1 line of equality.

Figure SI-2. Relationships between specific conductance (SC) in leach 3 $(H_2O_2+CO_2)$ and acid base accounting (ABA) parameters.

Figure SI-3. Total Ra and ²²⁸Ra/²²⁶Ra isotope ratios for the six rock types. Coal (n=3); Weathered Coal Refuse = W. Ref (n=14); Unleached Coal Refuse = U. Ref (n=20); Overburden = Overb (n=17); Shale (n=10); Pyrite (n=1). Box plots show median, 25% and 75% quartile ranges.

Table SI-4. Sample descriptions and associated PHREEQC input and output data for rapid leach samples: type 1 [element concentrations corrected for dilution of initial leach volume to 100 ml; concentration values original

TGS10B 004 refuseU 45.9 <0.05 -45.85 14700 27900 71.9 TGS11 004 refuseU 6.88 0.48 -6.4 2200 10600 31.3 TGS12 004 refuseU 280 <0.05 -279.95 89600 122000 264 TGS13 004 refuseU 4.06 93.5 89.44 1300 3940 1790 TGS14 004 refuseU 216 <0.05 -215.95 69000 97200 212 TGS15 004 refuseU 38.4 <0.05 -38.35 12300 21100 67.1 TGS17 004 refuseU 58.4 15.2 -43.2 18700 31800 318

Table SI-6. Chemical composition of rock samples used for rapid leach tests. letter AQN (increased conducted on a sched camples directed in HCLHNO3 (partial direction): MPA maxi

Analyte							MA_F MA_L1 MA_L3 MB_F MB_L1 MB_L3 SK_F SK_L1 SK_L3 MP_F MP_L1					
SC	Mean	8,110	2,350	6,060	8,240	2,180	8,170	21,100 5,480		25,900	11,400 6,270	
μ S/cm	Max	13,500 3,180		10,100	14,000 2,880		12,700	30,300 7,140		28,500		13,000 31,100
	Min	3,820	1,640	2,100	2,520	1,190	5,010	11,700 4,060		21,200	9,690	340
pH	Mean	3.6	4.1	2.3	3.6	4.4	2.1	2.2	2.8	1.3	6.8	4.6
	Max	5.8	6.2	3.3	$7.2\,$	6.8	2.5	2.3	2.9	1.5	7.6	7.6
	Min	2.5	2.6	1.8	2.7	2.9	$1.7\,$	2.0	2.7	$1.2\,$	6.3	$2.2\,$
Ca	Mean	NP	364	401	NP	295	463	NP	56.3	248.1	0.32	136
mg/L	Max	NP	446	501	NP	410	526	NP	56.3	248.1	0.41	491
	Min	NP	234	121	NP	42.5	398	NP	56.3	248.1	0.15	1.88
Mg	Mean	NP	30.1	67.9	NP	72.5	111	NP	27.5	36.3	168	58.6
mg/L	Max	NP	55.6	99.4	NP	144	197	NP	27.5	36.3	242	230
	Min	NP	15.0	16.9	NP	3.75	53.8	NP	27.5	36.3	35.0	0.313
$\rm Na$	Mean	NP	3.50	4.50	NP	51.0	82.1	NP	0.625	6.250	2,270	275
mg/L	Max	NP	5.00	9.38	NP	188	284	NP	0.625	6.250	2,660	731
	Min	NP	1.88	1.25	NP	1.88	1.25	NP	0.625	6.250	2,030	6.88
$\bf K$	Mean	NP	2.88	2.88	NP	4.69	9.27	NP	1.88	12.50	25.9	12.4
mg/L	Max	NP	5.00	6.88	NP	12.5	35.0	NP	1.88	12.50	54.6	29.2
	Min	NP	1.25	1.25	NP	1.25	1.25	NP	1.88	12.50	11.2	1.25
$\rm Fe$	Mean	969	190	971	1,160	49.3	1,390	9,525	863	5860	352	3,260
mg/L	Max	2,180	788	1,690	3,470	119	2,590	18,200 863		5860	590	32,300
	Min	20.9	5.25	66.3	11.4	0.125	466	1,850	863	5860	2.26	0.125
\mathbf{A} l	Mean	43.1	32.0	69.9	168	36.3	164	3,670	535	781	0.721	201
mg/L	Max	117	80.6	162	511	77.5	251	6,430	535	781	2.03	1,300
	Min	4.37	0.313	3.13	3.21	0.313	47.5	1,190	535	781	0.033	0.313
${\rm Mn}$	Mean	20.0	4.80	12.5	40.6	9.07	19.9	234	0.438	0.750	5.13	2.76
mg/L	Max	53.3	11.3	22.1	64.0	20.5	31.9	444	0.438	0.750	9.67	10.50
	Min	6.37	1.63	1.06	3.14	0.044	7.38	44.0	0.438	0.750	0.590	0.013
SO_4	Mean	5,040	1,220	2,680	8,450	1,230	4,130	57,300 2,850		17,100	4,920	4,890
mg/L	Max	9,840	1,870	4,310	26,300 1,730		6,690	97,600 2,850		17,100	8,130	52,400
	Min	1,820	675	969	1,670	488	1,960	20,900 2,850		17,100	1,860	3.13
HCO ₃	Mean	NP	2.27	0.006	NP	19.8	0.003	NP	0.007	0.001	NP	31.0
mg/L	Max	NP	9.81	0.022	NP	82.3	0.006	NP	0.007	0.001	NP	228
	Min	NP	0.003	0.001	NP	0.007	0.002	$\ensuremath{\mathbf{NP}}\xspace$	0.007	0.001	NP	0.003
Cl	Mean	NP	5.38	2.63	NP	7.63	4.00	NP	5.63	9.38	1,370	22.2
mg/L	Max	NP	11.9	5.00	NP	15.6	6.88	NP	5.63	9.38	2,020	101
	Min	NP	2.50	1.25	NP	5.00	1.88	NP	5.63	9.38	397	2.50

Table SI-10. Summary of field chemistry and leachate results for the 10 paired field sites. L1 (deionized water), and L3 (30% H2O2+10% CO2).

Table SI-11. Spearman rank correlation coefficient (r) matrix for XRD, acid-base account paramɾ [r-values multiplied by 100 and rounded; only values significant at α = 0.001 shown]

Table SI-6. Chemical composition of rock samples used for rapid leach tests.
[elemental analysis conducted on ashed samples digested in HCl+HNO3 (partial digestion); MPA, maximum potential acidity; NP, neutralization pote

1 - Cravotta and Brady (2015) Applied Geochemistry 62, 108-130

- 2 Sena et al. (2014) Water, Air and Soil Pollution 225, 1-14
3 Cravottta (2008) Applied Geochemistry 23, 166-202
n.p. not yet published
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Table SI-11. Spearman rank correlation coefficient (r) matrix for XRD, acid-base account parameters, and leachate chemistry $[{\rm r}\text{-}values$ multiplied by 100 and rounded; only values significant at α = 0.001 shown]

XRD-ABA-leachate correlation matrix. Significant Spearman (rank) correlation coefficients are shown for rock leaches 001, 002, and 003, combined (excluding blanks), in the first table and, then, separately for leaches 001 and 003. Each correlation matrix includes acid-base-accounting, XRD, and selected leachate chemistry parameters. For quantification of XRD, sulfur and carbonate minerals were assigned values of 2 (major), 1 (minor), or 0 (not identified).

The overall correlation matrix and charts support the hypothesis that samples containing sulfide and sulfate minerals (Sminl; FeSulfide; FeSulfate) have higher total sulfur content and corresponding potential for acid formation (expressed MPAppt = S% x 31.25) . Likewise, samples containing Calcite and Dolomite (Carbonate) have higher neutralization potential (NP) than other rock types/samples.

The samples containing sulfur minerals generate the highest conductivity and associated measures of salinity.

Also, the computed net alkalinity (negative of net acidity) for the leaches, combined, is correlated with the net neutralization potential (NP-MPA) and is positively correlated with samples containing carbonate minerals and negatively correlated with those containing FeSulfate minerals. The strongest correlations between ABA parameters and net alkalinity are indicated for leach 003.

The salinity (SPC, TDS, or osmotic pressure) of the leaches is not significantly correlated to the sum of MPA+NP (MPANPppt). The strongest predictor of salinity is the total sulfur (MPAppt) concentration and the presence of sulfur minerals (Sminl).

The correlations considering only leaches 001 or 003 support the hypothesis that leach 001 liberates sulfur and iron mainly from FeSulfate minerals (slightly higher coefficients than those for leach 003). Identification of FeSulfide does not seem to be particularly informative for predicting water chemistry. Generally, the significance of correlations between ABA parameters and salinity parameters increases for leach 003 (which seems to mobilize Ca from carbonates, whereas leach 001 mobilizes Ca from gypsum).

